

Ocean acoustic reverberation tomography

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Seismic wide-angle imaging using ship-towed acoustic sources and networks of ocean bottom seismographs is a common technique for exploring earth structure beneath the oceans. In these studies, the recorded data are dominated by acoustic waves propagating as reverberations in the water column. For surveys with a small receiver spacing (e.g., <10 km), the acoustic wave field densely samples properties of the water column over the width of the receiver array. A method, referred to as ocean acoustic reverberation tomography, is developed that uses the travel times of direct and reflected waves to image ocean acoustic structure. Reverberation tomography offers an alternative approach for determining the structure of the oceans and advancing the understanding of ocean heat content and mixing processes. The technique has the potential for revealing small-scale ocean thermal structure over the entire vertical height of the water column and along long survey profiles or across three-dimensional volumes of the ocean. For realistic experimental geometries and data noise levels, the method can produce images of ocean sound speed on a smaller scale than traditional acoustic tomography. © 2015 Acoustical Society of America.

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1. INTRODUCTION

While the oceans are principally vertically stratified in temperature and density, circulation and mixing processes produce significant secondary structure. Tides, winds, and buoyancy gradients (created by surface heat and freshwater fluxes) drive flow in the oceans. The energy behind this flow is transformed from larger patterns of circulation, including thermohaline circulation, gyres, and upwellings, into internal waves and dynamic current energy, such as eddies, cascading through a range of smaller scales leading into turbulence and molecular dissipation (e.g., Wunsch and Ferrari, 2004). These flows play a crucial role in driving diapycnal (across density strata) mixing (e.g., Munk, 1981; Polzin *et al.*, 1997; Garrett, 2003), which is a primary means for transferring heat energy vertically in the water column, and a key factor for models of ocean circulation. They also play important roles in redistributing nutrients and biomass and re-suspending sediment (e.g., Mann and Lazier, 2005; Pomar *et al.*, 2012). Mixing processes in the oceans are still poorly understood, motivating the development of new methods for measuring the internal structure of the oceans on small spatial scales and over large depth range and distance.

Acoustic energy propagates efficiently in the oceans, particularly at frequencies less than a few hundred hertz, and sound is now widely used for remote sensing of the ocean environment. Apart from pressure, which has a predictable influence on the speed of sound in the oceans, and salinity, which is important in some environments, temperature is the foremost factor affecting sound speed, and acoustic measurements allow large areas of the oceans' interior to be

thermally monitored by remote sensing. Ocean acoustic tomography is a technique that maps sound speeds across large regions of ocean by measuring the travel time of sound pulses transmitted and received between horizontally separated sources and receivers. In the presence of currents and other flows, small changes in temperature produce detectable changes in acoustic wave speed. Several successful experiments have been conducted using a variety of source-receiver geometries, most of which exploit properties of the sound propagation channel, avoiding bottom interactions. The measured travel times of acoustic energy include the accumulated effects of temperature variations along the acoustic path; hence, measurements by this technique are inherently averaging. The measurements also average variability over depth, since the ray paths of acoustic energy cycle vertically throughout the water column (see, for example, reviews by Dushaw *et al.*, 1993; Munk *et al.*, 1995; and Dushaw *et al.*, 2001).

In recent years, “seismic oceanography” studies have employed ship-towed hydrophone streamers and acoustic sources to image detailed ocean structure (e.g., Holbrook *et al.*, 2003; Nandi *et al.*, 2004; Holbrook and Fer, 2005; Nakamura *et al.*, 2006; Kraumann *et al.*, 2008). In these studies, an acoustic pulse is generated a few meters below the sea surface by a ship-towed compressed-air (or alternative) source. Impedance (the product of acoustic wave speed and density) contrasts in the water column produce reflections of the acoustic wave field that are recorded by the hydrophone streamer, which means that the recorded wave field is more sensitive to variations of the acoustic properties than to the acoustic properties themselves. As the vessel moves through the water, the system records a long vertical section of impedance boundaries within the water column. Marine reflection seismology is sensitive to thermohaline

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